

Sandstone geomorphology of the Golden Gate Highlands National Park, South Africa, in a global context

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Dates:

Received: 04 Jan. 2010
 Accepted: 12 Nov. 2010
 Published: 14 Mar. 2011

How to cite this article:

Grab, S.W., Goudie, A.S.,
 Viles, H.A. & Webb,
 N., 2011, 'Sandstone
 geomorphology of the
 Golden Gate Highlands
 National Park, South Africa,
 in a global context', *Koedoe*
 53(1), Art. #985, 14 pages.
 DOI: 10.4102/koedoe.v53i1.985

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The Golden Gate Highlands National Park (GGHNP) is well known for its impressive sandstone formations. While previous geoscience research in the park has focused on geology, palaeontology, slope forms and the prominent lichen weathering, remarkably little has been written on the diversity and possible origins of sandstone phenomena in the region. The objectives of this study were (1) to present a geomorphological map of prominent and interesting landforms for particular portions of the park and (2) to document the variety of macro- and microscale sandstone formations observed. During field work, we undertook global positioning system measurements to map landforms and, in addition, measured the dimensions of several landform types. A Schmidt hammer was used to conduct rock hardness tests at a variety of localities and lithologies for comparative purposes. We identified and mapped 27 macro- and microscale sandstone landforms, of which 17 are described in detail. It is demonstrated that for the most part, the landforms are a likely product of surface lithological reactions to a regional climate characterised by pronounced multitemporal temperature and moisture shifts, recently and in the past. However, many of the geomorphological processes producing landforms are controlled by microclimates set up by factors such as macro- and microtopography.

Conservation implications: The GGHNP is best known for its geological, geomorphological and palaeontological heritage. This paper highlights the diversity of sandstone geomorphological phenomena, many of them rare and 'unique' to the region. Not only are these landforms of aesthetic interest to tourists, but they also provide microhabitats for biota. Thus, conservation of biota requires associated conservation of geo-environments where they are established.

Introduction

The work *Sandstone landforms* (Young, Wray & Young 2009) provides a global view of our current understanding of macro- and microscale sandstone landforms and their associated geomorphic processes. It provides numerous examples of sandstone landforms from around the world and highlights the many important new studies that have appeared from South America, Australia, central Europe and China, in particular. Although the book identifies the Clarens area of the northeastern Free State as one of the world's 61 major sandstone landscapes, relatively little is known about the diversity of sandstone phenomena and their process origins in southern Africa. Despite sandstones not generally being as well known for their distinctive landforms as are limestones and granites (Robinson & Williams 1994), the Golden Gate Highlands National Park (GGHNP) hosts some of southern Africa's most spectacular scenery. The GGHNP has been identified as a geotourism hotspot, with the defining feature being the 'red, cream and yellow weathered sandstone of the Clarens Formation that is deeply eroded to form bluffs and cliffs' (Reimold, Whitfield & Wallmach 2005:55). In recent years several World Heritage sites have been proclaimed in South Africa, and more than half of those currently listed and proposed represent entirely natural geological heritage (Reimold *et al.* 2005). The GGHNP falls within the Drakensberg World Heritage region, which boasts mixed heritage status owing to the presence of both culturally and naturally significant features. It is thus not surprising that the GGHNP has become a world-famous geotourism destination, with particular attractions including the Brandwag 'Golden Gate' (Figure 1), San rock art and vertebrate fossils.

The GGHNP was established in 1963 for the purpose of protecting a pristine area, but more specifically, to conserve the impressive sandstone formations (hence its name) and the montane and Afro-Alpine grassland biome (SANParks 2004). Previous earth science research in the park has focused mainly on the geological and palaeontological heritage and macroscale geomorphological evolution (e.g. Groenewald 1986; Kitching & Raath 1984; Spies 1969), slope forms and processes



FIGURE 1: The Brandwag 'Golden Gate', displaying Clarens Formation sandstone cliffs.

(Moon 1991; Moon & Munro-Perry 1988; Nicol 1973), biogenic weathering (Büdel *et al.* 2004; Wessels & Büdel 1995; Wessels & Schoeman 1988; Wessels & Wessels 1991, 1995; Wessels *et al.* 1995), weathering basins (Cooks & Pretorius 1987) and the origins of 'cave' formations (Le Roux 1978). Despite these previous studies and the recognition of the regional geoheritage, the diversity of both macro- and microscale sandstone landforms has not yet been comprehensively documented or mapped in the GGHNP. To this end, our objective was, firstly, to provide a map of prominent and interesting sandstone phenomena for particular portions of the park, thus contributing to enhanced knowledge on geodiversity. Secondly, we aimed to document the variety of macro- and microscale sandstone morphologies (Table 1) and their possible process origins within a global context. Although we concentrate here on sandstone features, some of our observations may also be relevant to granitic terrains, which often have similar assemblages of weathering forms (see Twidale & Vidal Romaní 2005).

Study area

The GGHNP is located in the eastern Free State (28°31'S, 28°37'E) and covers an area of 120 km² across the Rooiberg Mountain Range (1892–2837 m a.s.l.), which forms part of the greater Drakensberg–Maluti mountain system (Figure 2). Although annual precipitation varies considerably across the region, it averages approximately 764 mm, of which most falls between November and April (Groenewald 1986). Summers are mild to warm (mean temperature ranges from 13 °C to 26 °C) and winters cold (mean temperature ranges from 1 °C to 15 °C). (Cooks & Pretorius 1987). Frost is widespread during the winter months and snow occasionally falls on the higher peaks in the park. The vegetation is typically Highlands Sourveld and *Themeda–Festuca* Alpine Veld, with identified plant communities representative of the *Tristachya leucothrix–Trachypogon spicatus* Grassland of moist mountain slopes and plateaus, and to some extent the *Merxmuellera drakensbergensis–Festuca caprina* high-altitude Austro-Afro Alpine Grassland

(Du Preez & Bredenkamp 1991; Kay, Bredenkamp & Theron 1993). The area is located on the African Surface of Partridge and Maud (1987) and its drainage systems have undergone subsequent incision (Moon & Munro-Perry 1988). The observed landscape has been influenced by ever-changing climates, particularly during the Quaternary period, which encountered an intensification of cold climate processes and also major changes in vegetation cover (Henderson *et al.* 2006). The landforms of the area thus have a long and complex geomorphic history.

Methodology

Large portions of the sandstone sections of the GGHNP were traversed, with sandstone phenomena photographed and recorded using a global positioning system device. The recorded phenomena could then be incorporated onto existing topographic and geological maps. For some landforms and topographic settings we undertook morphometric and rock hardness measurements using a tape measure and a Schmidt hammer (classic N-type), respectively. The Schmidt hammer indicates rebound values (*R*-values ranging from 0 to 100), which are proportional to the hardness of the rock. Rock hardness as described here is based on the indentation produced when pressure is exerted on the rock mass (as opposed to individual mineral hardness, which indicates resistance to scratching). Rock surfaces were first cleaned (smoothed) using a Carborundum, after which the hammer was applied to the smoothed surface. Mean rebound values were calculated from 30–50 readings per site. The objective with this was primarily to compare the hardness of Clarens Formation sandstone with that of other geological units in the park and consequently its potential strength variability in various spatial and geological contexts.

Young *et al.* (2009) classify sandstone phenomena into four hierarchical groupings, based on morphology rather than process origins. We classify the sandstone landforms observed in the GGHNP according to size and morphological attributes. However, given that microsurface forms such as case hardening and spalling (flaking) contribute to the development of other sandstone features (e.g. rock basins and tafoni), these are dealt with first as part of the geological structure which invariably impacts on landform genesis.

While a variety of landform types are discussed within separate subsections for ease of presentation, it should be acknowledged that several varieties of sandstone phenomena may be closely linked and have common developmental origins despite their different morphological surface expressions. Such differences may be due to macro- and microtopographic position, existing surface morphology and differences in underlying geological controls. For instance, rock doughnuts, rock basins and rock pans within a region may have common process geomorphic origins but the outcomes may spatially take on rather different visible forms.

Geological context

The sedimentary portion of the Late Carboniferous – Middle Jurassic Karoo Basin of southern Africa includes glacial

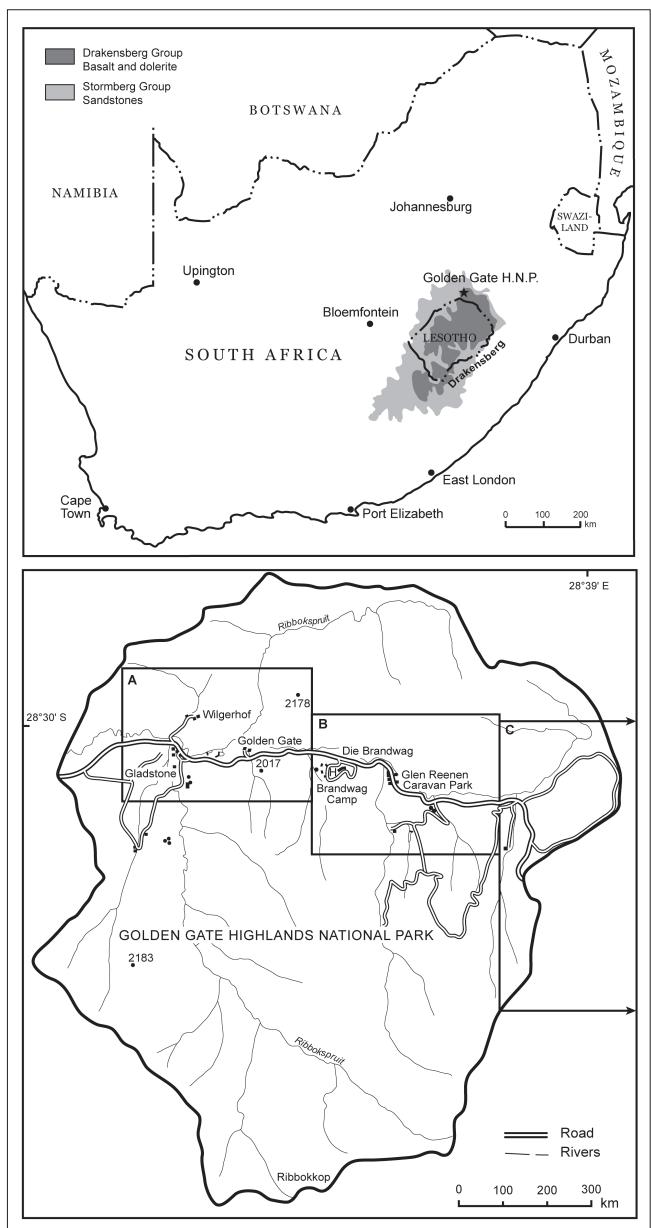


FIGURE 2: Location of the Golden Gate Highlands National Park and map of the park showing the insets for the geomorphological maps in Figure 3.

(Dwyka Group), marine, swamp and deltaic (Ecca Group), fluvial (Beaufort Group and parts of the Stormberg Group) and aeolian (upper part of the Stormberg Group) deposits, and is capped by the volcanic Drakensberg Group (Catuneanu, Hancox & Rubidge 1998; Catuneanu *et al.* 2002; Smith *et al.* 1998). Sills and dykes were emplaced in sedimentary strata around 183 Ma (Holzförster 2007) and in places may form up to 70% of the basin volume (Rowse & De Swardt 1976; Svensen *et al.* 2006). The Stormberg Group is made up of the Molteno, Elliot and Clarens Formations, all of which are represented in the GGHNP (Eriksson 1981, 1984, 1985; Groenewald 1986). While the Elliot Formation has a distinct lower contact with the Molteno Formation, it has a more gradual upper contact with the Clarens Formation (Bordy, Hancox & Rubidge 2004; Groenewald 1986). The succession reflects changing palaeo-environmental conditions (climatic and basin foreland tectonics) during deposition; from

dominantly fluvial (Molteno Formation) to increasingly aeolian (Clarens Formation) (Bordy *et al.* 2004; Eriksson 1981, 1984, 1985; Groenewald 1986).

Sandstones are produced through the cementation or lithification of sand grains through processes of pressure melting and recrystallisation, compaction and consolidation of incorporated clay minerals, and/or by the precipitation of authigenic minerals from water migration through pores (Robinson & Williams 1994). Although a wide variety of sandstones occur in the GGHNP, primarily owing to a diversity of depositional mechanisms, they are, for the most part, a product of compaction and consolidation over time. The Molteno Formation varies in thickness from less than 10 m to more than 100 m and consists of light-coloured sandstones of which the grains can range from fine to very coarse (Eriksson 1984). The Molteno Formation sandstones are primarily quartz-rich feldspathic wackes, have a high interstitial content and sometimes tend towards subordinate argillaceous and conglomeritic structures (Eriksson 1984). The Elliot Formation varies in thickness from approximately 28 m to more than 150 m and consists of red-coloured, argillaceous sandstone with grains ranging from fine to coarse (Eriksson 1985). The Elliot Formation is particularly well known in the GGHNP for its vertebrate fossil content (e.g. Groenewald 1986; Kitching & Raath 1984). The Clarens Formation is typically 115 m – 195 m thick and consists of light-coloured, fine-grained sandstones, sandy siltstones and mudstones, and subordinate coarser-grained components (Eriksson 1981). Fresh faces of the rock collected from boreholes and quarries are white, yet on exposure the rock weathers to give its characteristic tan or golden colour. However, there have been no detailed studies on the chemical and mechanical weathering processes causing such colour alterations. The sedimentary strata, which are generally approximately horizontally bedded, have been intruded by several fine- to coarse-grained dolerite dykes and a few sills in the GGHNP (Groenewald 1986). Some major faults occur in association with these intrusions and these are likely to have played a role in the genesis of the landforms, as has the presence of the bedding planes and joints within the sedimentary rocks.

Schmidt hammer testing of the rocks in the area provides estimates of the hardness characteristics of the various rock units (Table 2). The mean Schmidt hammer hardness *R*-value for the Clarens Formation sandstone, based on 30–50 readings from each of nine sites, was 48.91. This is marginally higher than the *R*-values that have been recorded at two other major sandstone geomorphological sites: approximately 45 for the Navajo sandstone of Utah (Irwin *et al.* 2009) and 41–44.7 for the Cambrian–Ordovician sandstones of Wadi Rum in Jordan (Goudie *et al.* 2002). The *R*-value for a dolerite dyke piercing the Clarens Formation sandstone was 63.0, while that for Drakensberg Group basalt was 55.92. Sandstone of the Elliot Formation had an *R*-value of 58.12. Thus, all the main rock units in the area are relatively hard. However, the wide variety of rock hardness values (compressive strength) obtained across sandstone surfaces in the GGHNP attests to both macro- and microscale variability in porosity,

TABLE 1: A scale hierarchy of sandstone landforms in the Golden Gate Highlands National Park.

Large (> 100 m)	Medium (10 m – 100 m)	Small (< 1 m)
• Cliffs	• Pillars • Butresses • Arches • Rock shelters • Alcoves • Caves • Tafoni	• Honeycombs
• Interfluves Plateaus	• Pedestal rocks • Rock basins • Rock doughnuts	• Tessellations • Polygonal boxwork • Fluting • Gutters • Trittkarren • Lichen pits
• Low-angle convex valley side slopes	• Turf slides • Boulder fields	• Tessellations • Polygonal boxwork • Fluting • Gutters • Trittkarren • Lichen pits

composition and bonding of mineral grains, and the extent of surface weathering or case hardening. Although not yet locally tested, it is thus likely that many of the sandstone morphologies at the GGHNP are a product of differential rates and processes of weathering and erosion across microscales, owing at least in part to rock hardness variability.

'Blueberries' (concretions)

Calcareous and ferruginous concretions are particularly common to the sandstones of the Clarens Formation (Groenewald 1986), where they impact on the fine-scale spatial dynamics of weathering processes. Such concretions are sometimes informally referred to as 'blueberries', possibly owing to their nodular spheroidal structure (Squyres & Knoll 2005). According to Mahaney *et al.* (2008), it is thought that such surface nodules are the product of subsurface fluid migration through porous sandstone. Fluids become enriched with iron and calcite, with precipitates eventually forming around a central nucleus of fine sandstone particles (Mahaney *et al.* 2008). These are released when the host sandstone is weathered and produce a surface lag.

Case hardening

An important and common attribute of many sandstone surfaces is the development of a surface crust or 'outer [shell] more resistant to erosion than interior material' (Dorn

2004:118). Crust development is primarily caused by the deposition of solutes when pore water migrates towards the surface and evaporates (Robinson & Williams 1994). However, the processes causing case hardening may be polygenetic in origin and vary between sites. For example, it has been found that micro-organic biofilms may aid the cementation process to produce case-hardened surfaces in Jordanian sandstones (Viles & Goudie 2004). Examples of case hardening occur throughout the GGHNP where the iron-rich crusts produce hardened, reddish surfaces. The importance of case hardening cannot be underestimated: on the one hand, it protects underlying rock from weathering and, on the other, it reduces localised infiltration and so facilitates enhanced surface runoff. Both these functions consequently impact on weathering and erosion on the surrounding uncrusted rock surfaces. In addition, the disintegration of the surface crust provides new irregularities that may promote localised weathering and erosion and consequently lead to new phenomena such as rock basins (weathering pits) (Goudie & Migón 1997).

Surface spalling (exfoliation, flaking)

In arid and semi-arid regions (or microsettings) with seasonal rainfall (humidity) patterns, sandstone may be susceptible to spalling or flaking (i.e. thin rock flakes peeling off). In such instances cracks develop parallel to the surface, typically at a depth of between 5 mm and 50 mm (Robinson & Williams 1987, 1992). Spalling of rock surfaces is common where cracks develop below a case-hardened surface, or along the backwalls of tafoni caverns and rock shelters. It is understood that a primary mechanism for such sandstone spalling is the dissolution of silica and other cementing agents during wetting cycles and the translocation in the surface layers during drying cycles (Robinson & Williams 1992). Crystallisation of salt during desiccation cycles has also been observed to cause spalling (Goudie & Viles 1997; Holmer 1998). Observations on the backwalls of cavernous weathering features show that in the GGHNP there are salt efflorescences associated with spalling rock. It has also been suggested that contemporary sub-zero air temperatures at the GGHNP during the winter period may contribute to frost action (spalling) where moisture is locally available (Marker 1990); however, freeze-thaw processes are considered contentious given that rock thermal data required to prove such a mechanism have not been forthcoming (Hall, Marker & Le Roux 1991).

TABLE 2: Hardness values as measured with a Schmidt hammer (classic N) for rock types in the Golden Gate Highlands National Park.

Rock type	Site characteristics	Mean R-value (corrected)	Sample size
Clarens sandstone	Horizontal open surface	57.90	50
Clarens sandstone	Horizontal surface of boulder	52.71	30
Clarens sandstone	Horizontal surface of boulder	47.96	30
Clarens sandstone	Horizontal open, dry surface	51.15	50
Clarens sandstone	Horizontal open surface after wetting	49.95	50
Clarens sandstone (case hardened)	Rock shelter caprock vertical face	45.32	50
Clarens sandstone (weathered)	Vertical backwall of rock shelter	46.78	50
Clarens sandstone (iron crusted)	Horizontal open surface	45.02	50
Clarens sandstone (baked)	Vertical surface on ridge adjacent to dolerite dyke	43.42	50
Dolerite	Vertical face of dolerite dyke	63.20	50
Basalt	Vertical face of basalt outcrop	55.92	50
Elliot sandstone	Vertical face	58.12	50

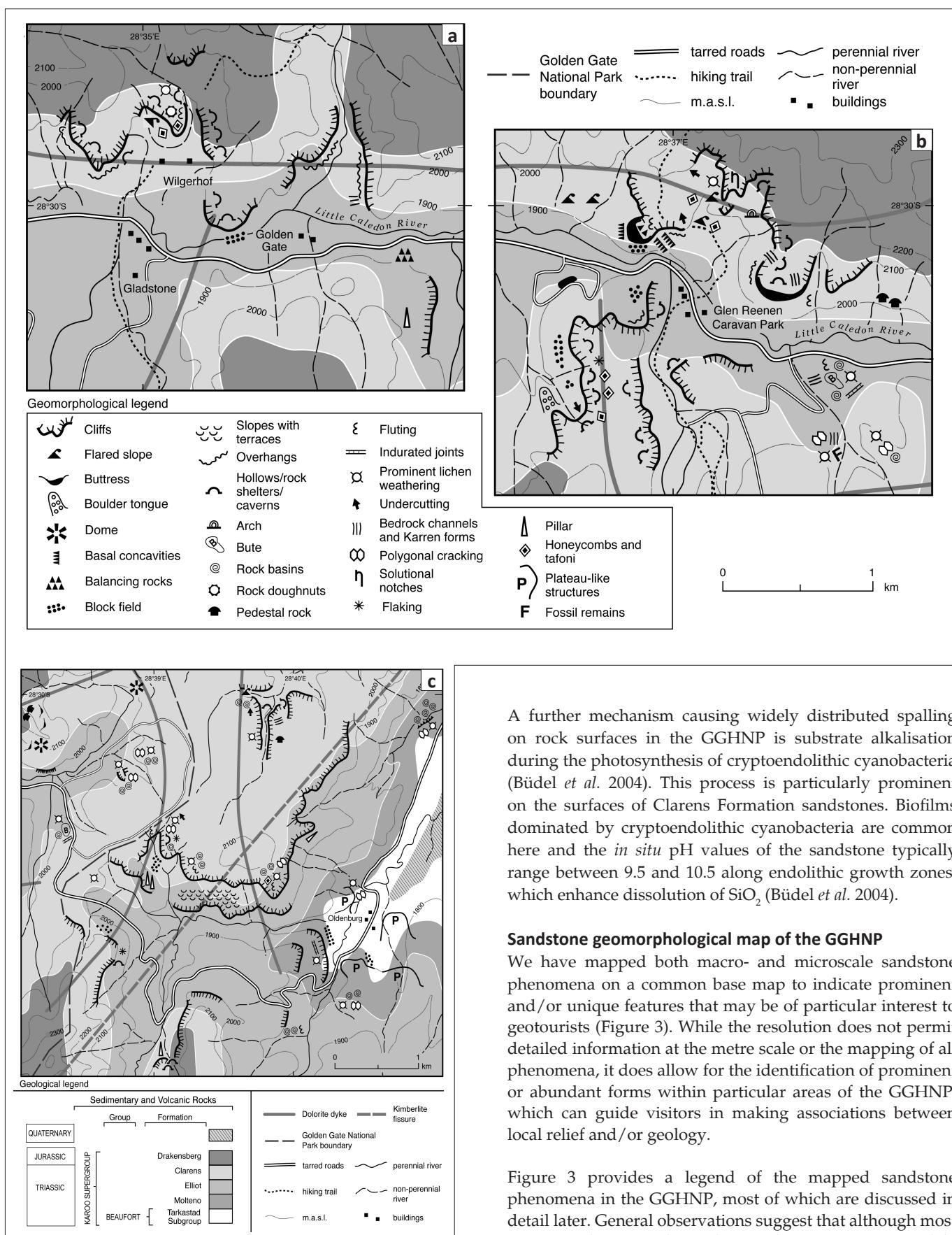


FIGURE 3: Geomorphological map of sandstone phenomena in the Golden Gate Highlands National Park (geological overlays modified after Geological Series 2828 Harrismith, Council for Geoscience, 1998; localities of subfigures a–c as shown in Figure 2).

A further mechanism causing widely distributed spalling on rock surfaces in the GGHNP is substrate alkalinisation during the photosynthesis of cryptoendolithic cyanobacteria (Büdel *et al.* 2004). This process is particularly prominent on the surfaces of Clarens Formation sandstones. Biofilms dominated by cryptoendolithic cyanobacteria are common here and the *in situ* pH values of the sandstone typically range between 9.5 and 10.5 along endolithic growth zones, which enhance dissolution of SiO_2 (Büdel *et al.* 2004).

Sandstone geomorphological map of the GGHNP

We have mapped both macro- and microscale sandstone phenomena on a common base map to indicate prominent and/or unique features that may be of particular interest to geotourists (Figure 3). While the resolution does not permit detailed information at the metre scale or the mapping of all phenomena, it does allow for the identification of prominent or abundant forms within particular areas of the GGHNP, which can guide visitors in making associations between local relief and/or geology.

Figure 3 provides a legend of the mapped sandstone phenomena in the GGHNP, most of which are discussed in detail later. General observations suggest that although most macro- and microscale sandstone phenomena occur widely within the various sedimentary formations of the Karoo Supergroup, the uppermost broad sandstone plateaus of the Clarens Formation host the highest density and greatest

variety of sandstone landforms. Notably, such plateaus are topographic positions capable of storing considerable quantities of surface water during the wet season (particularly where infiltration is relatively low or drainage is poor), while also being subjected to substantial heating and cooling cycles because of exposure. To this end, Paradise (2002) found that aspect-controlled weathering processes and associated landform production are particularly pronounced in regions where wetting-drying and heating-cooling cycles are important weathering mechanisms, such as have been observed in Petra, Jordan. Dolerite dyke intrusions are also likely to have influenced landform development within the adjacent sandstones in the GGHNP, and we note phenomena such as flared slopes, a rock arch, caverns, honeycombs and rock basins adjacent to some of the dykes.

Slope forms

1. General slope forms and processes: As a consequence of fluvial incision by the Little Caledon and its tributaries, the GGHNP has a variety of slope forms, ranging from horizontal sandstone pavements along broad interfluves to vertical cliffs along valley margins. The variable nature of bedrock forms (gradients) has to a large extent controlled the development of microscale sandstone phenomena through controlling

the amount of solar radiation received at the rock surface (rock surface microclimatology), moisture retention, surface water movement and gravity. There is likely feedback as the microscale phenomena contribute to the further evolution of the larger slope forms.

The prominent slope forms in the GGHNP are rectilinear cliff-talus slope combinations and bedrock slopes inclined at $30^\circ - 35^\circ$ (Moon & Munro-Perry 1988). The presence of a resistant and relatively massive caprock (normally Clarens Formation sandstone) permits parallel slope retreat and the preservation of rectilinear slopes where the rock mass strength values are higher than 70. However, where values are lower, the rock is too weak to preserve the Richter slope forms and thus undergoes a process of 'slope replacement' (Moon 1991; Moon & Munro-Perry 1988). Most valley sides typically have near-vertical cliffs and/or steeply inclined ($> 35^\circ$) bedrock surfaces, below which is a zone of discontinuous debris cover ('vegetated islands') on somewhat gentler slopes ($20^\circ - 35^\circ$). These 'vegetated islands' (Figure 4a) appear to be the result of peat rafts' sliding over the underlying bedrock, although there is evidence that the movements are complex and involve diffusive, waterborne erosion of sediment as well as sliding. Finally, there is a zone of debris accumulation on lower slope units approaching the valley floor (Moon &

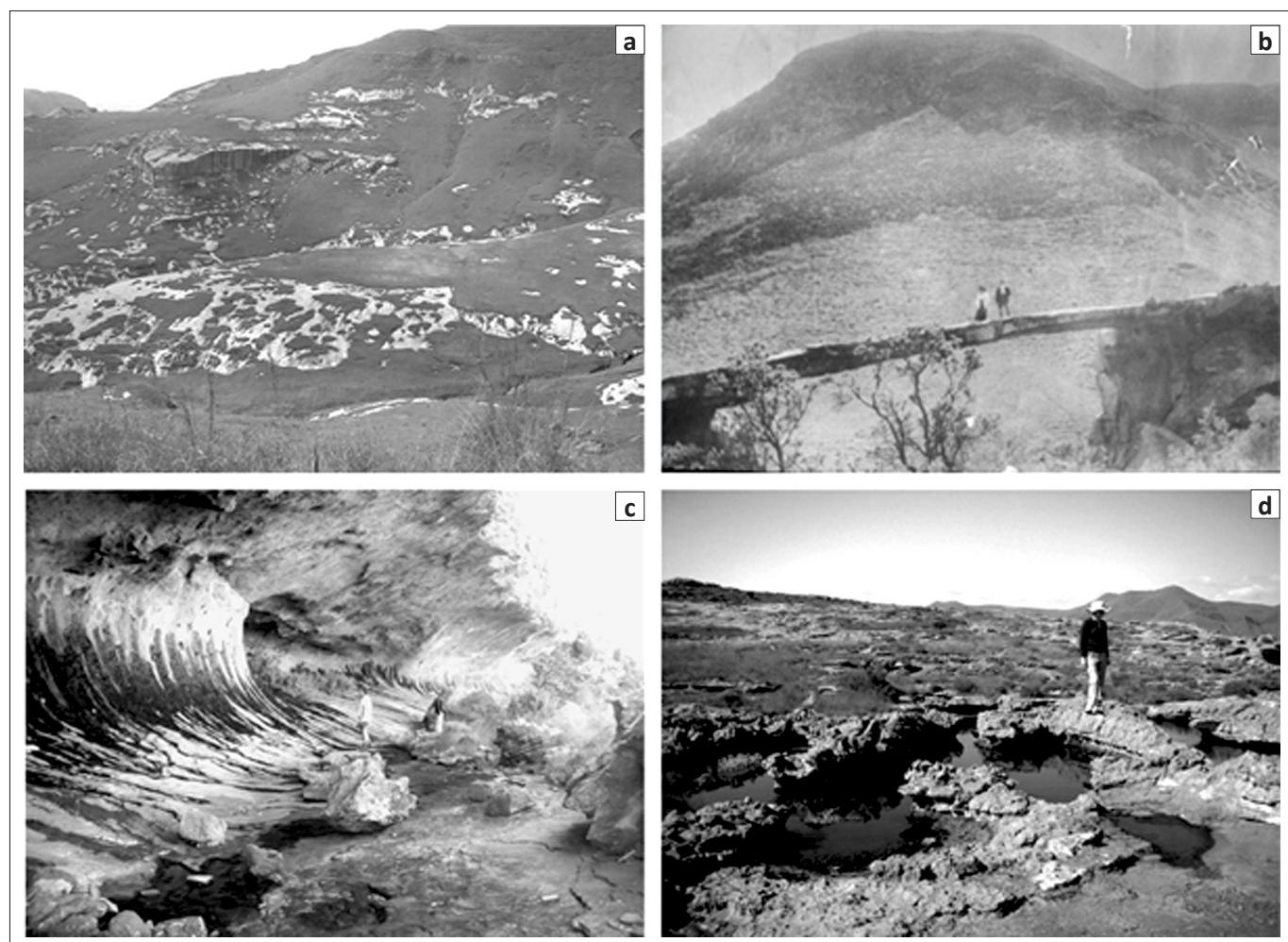


FIGURE 4: (a) 'Vegetated islands' overlying sandstone outcrops, (b) a natural stone bridge spanning the Zulu Hoek stream (since collapsed; electronic copy of original photo provided by Johan Taljaard, GGHNP), (c) water seepage and associated black staining at the rear of a cavern and (d) large rock basins with distinct water level notches.

Munro-Perry 1988; Nicol 1973). The number and size of some of the blocks on slopes below the cliffs is impressive and shows the importance of rock falls from the cliffs above.

Prominent 'cirque-shaped' hollows frequent several of the south-facing slopes in the GGHNP and these were initially thought to have originated from nival and frost-related processes during past Pleistocene cold phases, as were rockfall debris and thick colluvial mantles along the valley margins (Marker 1989; Nicol 1973). However, such cold-process origins were subsequently contested (Le Roux & Marker 1990) and have not been given much further consideration.

2. Cliffs and pillars: Cliffs and pillars represent steep bedrock exposures, especially along deeply incised valleys or around the periphery of plateaus (Robinson & Williams 1994). The conspicuous cliffs of the GGHNP are most prominent along the Caledon Valley and are primarily associated with the massive and strongly cemented fine-grained Clarens Formation sandstones. Although cliffs are also typical of the Molteno and Elliot Formations, the generally thicker and more stable Clarens Formation displays the most impressive and extensive cliff development in the region. The height of the cliffs, which in places exceed 100 m to make the cliffs among the highest in the world (Moon & Munro-Perry 1988), is to some extent controlled by the thickness of the sandstone formations (i.e. sedimentary beds) and the associated extent of basal weathering and erosion.

The best-known cliff in the GGHNP is the 'Brandwag buttress' (Figure 1), alongside which the Clarens Formation sandstone has been intruded by a dolerite dyke, and consequently has yielded a hard quartzitic sandstone with greater resistance to weathering (Groenewald 1986). The recent collapse (i.e. during the last few hundred years) of joint-bounded blocks at a few localities also provides evidence that cliff development is an ongoing, contemporary process in the GGHNP. The formation of pillars is limited in the GGHNP as joints are relatively widely spaced and the sandstone structures are massive and coherent. Nevertheless, block structures along cliffs have developed in places owing to accelerated weathering and erosion along more closely spaced joints. Another feature of the valley side slopes is the local development of summit and buttress tors.

Arches (natural stone bridges)

Despite their unusual and aesthetic appeal, natural arches and bridges are among the least studied of sandstone landforms worldwide. Fairbridge (1968) classified features into 'arches' (which do not span a stream channel) and 'bridges' (which span a stream channel). A photograph (Figure 4b) and painting by Johan van Reenen, dating back to between 1930 and 1940, depict the former occurrence of a natural stone bridge spanning the Zulu Hoek stream in the park. Although few known natural stone bridges exist in the region today, some small natural arches presently occur in the park and the surrounding area. Elsewhere, such arches are said to be the product of enhanced weathering and water infiltration along

parallel joints on cliff faces, eventually leading to material collapse and production of the 'window area' (Leisure, Kind & Busby 2007).

Pedestal rocks (mushroom-shaped pillars)

Relatively little research has focused on sandstone rock pedestals and mushroom-shaped pillars, also described elsewhere as rock columns and bulbous pinnacles (Mueller & Twidale 2002). Nevertheless, it is understood that such forms are the likely result of differential rates of weathering, which are enhanced along existing fracture-defined blocks and sheets in bedrock. Pedestals occasionally occur on slopes in the GGHNP; these were most likely initiated beneath a former mantle of regolith, which has subsequently been eroded and left behind such bedrock landforms of etch origin, as suggested by Mueller and Twidale (2002) for similar landforms identified in New Mexico, USA. Pedestals often have a characteristic basal concavity, above which are broader bedrock cappings, resembling a mushroom (Twidale & Campbell 1992). Although such bedrock forms may have been initiated through original subsurface etching processes, once they have been exhumed, differential rates of weathering and erosion would have continued along existing joints and areas of locally weaker rock strength, particularly from induced fracturing (Ollier 1978). The formation of these forms may also have been controlled by spatial dynamics such as distance from the ground surface and aspect.

Cavernous weathering

Cavernous weathering has been described as 'differential weathering of a rock surface' which represents 'areas of material loss surrounded by intervening areas of stable rock surface (with minimal material loss)' (Turkington & Phillips 2004:665). International research on cavernous weathering has focused on both microcavernous forms (e.g. negative spheroidal weathering forms [Wagner 1913], dewholes [Demek 1964] and honeycombs [e.g. McBride & Picard 2004; Rodriguez-Navarro, Doehne & Sebastian 1999]), and macrocavernous forms such as tafoni (e.g. Martini 1978; Matsukura & Tanaka 2000; Mottershead & Pye 1994), alcoves (Lamb *et al.* 2006) and rock overhangs (locally also known as rock shelters or caves). To this end, there has been limited discussion concerning the likely successive evolution of forms from honeycombs (embryonic tafoni) to tafoni to so-called rock shelters ('mature tafoni' or 'tafoni caverns' as described by Mellor, Short & Kirkby 1997). Cavernous weathering is particularly prominent in crystalline rocks with a medium to coarse grain structure, including granites and sandstones (Jennings 1968; Mellor *et al.* 1997), and is thus active within the sandstones of the GGHNP.

1. Honeycombs: Honeycombs are the smallest of the cavernous weathering types (only few tens of centimetres across), consisting of closely spaced pits which are bounded by thin raised lips or walls (McBride & Picard 2004). Such weathering of microforms have previously been described as stone lattices, stone lace or fretting (Mustoe 1982), but are more commonly also known as alveoli (e.g. Benito, Machado & Sancho 1993; Grisez 1960). In some instances honeycombs may develop within larger cavernous forms such as tafoni,

but are said to be affected more strongly by lithological heterogeneities (Martini 1978). In fact, Groenewald (1986) observed honeycombs in the GGHNP and similarly ascribed their formation to differential weathering associated with calcareous concretions. International research suggests that honeycombs are typically a product of solutional processes and the crystallisation of salt in the pits (e.g. Lambert 1980; McBride & Picard 2004; Young 1987), while case hardening by iron oxide or calcium carbonate offers greater resistance to weathering along the surrounding, less denuded walls (Campbell 1999; McBride & Picard 2004). The spatial occurrence of honeycombs in the GGHNP is likely controlled by local sandstone permeability, the extent of sandstone cementation and mineralogy.

2. Tafoni, caverns, alcoves (rock shelters, caves): Tafoni are relatively large (several metres across) cavernous weathering forms with overhanging hoods, arch-shaped entrances, concave inner walls, overhanging margins and relatively smooth, gently sloping, debris-covered floors (McBride & Picard 2004; Mellor *et al.* 1997). The formation of tafoni has been ascribed to differential weathering, granular disintegration and flaking associated with microlithological and microclimatological differences along rock walls. Meteoric water migration through megapores and associated dissolution and reprecipitation of calcareous and saline cements, crystallisation pressure, hydrolysis, hydration and/or diffusion, and wind deflation have also been noted as particular contributors (Cooke, Warren & Goudie 1993; Martini 1978; Matsukura & Tanaka 2000; McBride & Picard 2000; Mellor *et al.* 1997; Mottershead & Pye 1994). However, water-based weathering processes may be temporally and spatially constrained at some tafoni sites, such as those reported from arid regions of Antarctica, where thermal stress may take on a more important weathering function (Hall & André 2006).

The GGHNP hosts a variety of variably sized tafoni forms (ranging from only a few to several tens of cubic metres), which most likely represent various stages of tafoni evolution, as described by Mellor *et al.* (1997). It would appear that several of the elongated rock shelters (caves) observed in the GGHNP (primarily in Clarens Formation sandstone) formed through the amalgamation of tafoni caverns that are in places separated by thin rock walls. Multiple vertical zones of tafoni and cavern development may occur within rock walls and are likely controlled by zonal lithological (rock hardness) and moisture differences (Matsukura & Tanaka 2000). The caverns are thought to enlarge in an upward and backward direction, thus potentially capturing subsurface water and causing water flow down the backwall (Mellor *et al.* 1997), as is frequently observed at sites in the GGHNP. As the tafoni enlarge over time, so the microclimate within the cavern is altered (primarily through greater contrasts in solar radiation levels across rock surfaces), thus influencing rock characteristics such as moisture and temperature. Consequently, different weathering processes may dominate on different parts of tafoni, such as case hardening on outer roof surfaces and granular disintegration on inner cavern surfaces (Mellor *et al.* 1997).

Another type of cavern occasionally observed in the GGHNP region is the alcove. An alcove has been defined as 'a steep-sided, arcuate cavity on the flank of a rock outcrop, typically produced by water erosion processes such as solution and spring sapping or seepage' (Nash 2000:14). Thus, several of the so-called caverns or caves in the GGHNP, and possibly even large tafoni which owe their origin to seepage and dissolution processes, might be classified as alcoves. Water is often found to be seeping out from the rear of these features, producing black staining on the backwalls (Figure 4c). However, as Lamb *et al.* (2006) point out, the origin of alcoves remains unclear given that the ability of seepage to erode bedrock is not yet comprehensively understood, and also because other mechanisms might have contributed to such arcuate cavities.

According to Le Roux (1978), caves in the GGHNP are a consequence of exudation, whereby weathering processes are accelerated at the base of cliffs, in part owing to the presence of shale or clay-bearing sandstone layers. Many of the caverns observed in the GGHNP are located at the lower boundary of the Clarens Formation where it merges with the siltstones and shale beds of the Elliot Formation. Here we observed that brittle fractures are common phenomena, possibly not so much due to weathering in the clay-like rocks, but rather due to mass failure of bedrock below the recess and beyond the dripline of water coming over the cliff (see also Young *et al.* 2009).

Rock basins (weathering or solutional pits, gnammas)

Rock basins are relatively circular or oval and usually flat-floored and steep-sided bedrock depressions ranging from a few centimetres to several metres in diameter and depth (Domínguez-Villar *et al.* 2009) (Figure 4d). Rock basins have also been referred to as gnammas (Domínguez-Villar *et al.* 2009), weathering pits (Alexandrowicz 1989; Domínguez-Villar 2006; Goudie & Migón 1997), potholes (Chan *et al.* 2005; Graham & Wirth 2008; Johnsson 1988), solution pans (Dzulynski & Kotarba 1979), *Opferkessel* (Hedges 1969), pias (Vidal Romaní 1989) and, locally, as rock tanks (King 1942) or weathering basins (Cooks & Pretorius 1987). Rock basins may contain standing water for short durations after rainfall or for several weeks to months during wet periods or seasons. Rock basins, which most typically develop in granite and sandstone, are understood to be polygenetic in origin (Domínguez-Villar *et al.* 2009; Dzulynski & Kotarba 1979). Processes include chemical weathering associated with dissolution (Domínguez-Villar *et al.* 2009; Twidale 1982; Wray 1997) and mechanical erosion associated with, for instance, convective (circulatory) currents within the standing water (Dzulynski & Kotarba 1979), wash and deflation (Alexandrowicz 1989; Graham & Wirth 2008), undercutting of the rock basin sides (Goudie & Migón 1997), and the plucking action of lichen thalli during the drying phase (Moses & Smith 1993). Complex biofilms may accumulate at the base of rock basins; these are known to dissolve the cement between sandstone grains and also act as biological sealants to water infiltration (Chan *et al.* 2005). Cooks and Pretorius (1987) have identified blue-green algae below the rock basins' seasonal water lines in the GGHNP and suggested that associated biochemical activity

would enhance weathering there. Solutional weathering is prominent along joints and bedding planes and may cause intense etching, particularly in quartz-rich sandstone (Young *et al.* 2009). Cooks and Pretorius (1987) similarly observed rock basin development in the GGHNP at sites where joints intersect at the surface. In addition, some of the pans have well-developed water level notches which are probably the result of solutional attack.

Rock basins in the GGHNP are commonly 15 cm – 150 cm wide and approximately 5 cm – 30 cm deep. However, Cooks and Pretorius (1987) report features up to 16 m in length and 1 m in depth. Examples of isolated deep pits (> 1 m deep) and exceptionally wide (several metres in diameter), shallow bedrock basins (also referred to locally as 'tarns') occur elsewhere on Clarens Formation sandstones, such as in the Sehlabathebe National Park, southern Drakensberg. In some cases spillways develop at the lowest point of the basin rim where erosion eventually produces a confined water exit route, below which drainage microchannels usually develop. The overflow of water during summer is an effective mechanism to remove weathered products accumulating in the rock basins. Less frequently, a staircase of rock basins with connected drainage microchannels may develop down

sandstone surfaces, while in other cases closely spaced basins may eventually coalesce. Rock basins may also be associated with incision into polygonally cracked surfaces (Robinson & Williams 1992) or with raised conical forms such as rock doughnuts (Netoff & Shroba 2001; Young *et al.* 2009).

Rock doughnuts

Rock doughnuts were first described by Blank (1951:822), who described them as 'weather pits of various sizes on granite' which are 'surrounded by raised annular rims'. Although rock doughnuts are relatively rare phenomena which have received little research attention, they have also been reported from a variety of sandstone terrains in Australia (Twidale & Campbell 1998; Young *et al.* 2009) and from south-eastern Utah, USA (Netoff & Shroba 2001). Rock doughnuts belong to a group of conical-shaped raised landforms with a central summit 'pit' or rock basin, examples of which were found at one locality in the GGHNP (Figures 3 and 5a). Rock doughnuts ranging between 5 cm and several metres high have been reported in Australia and the USA and have pit diameters ranging from approximately 20 cm to 10 m (Netoff & Shroba 2001; Twidale & Campbell 1998; Young *et al.* 2009). The ones in the GGHNP are typically

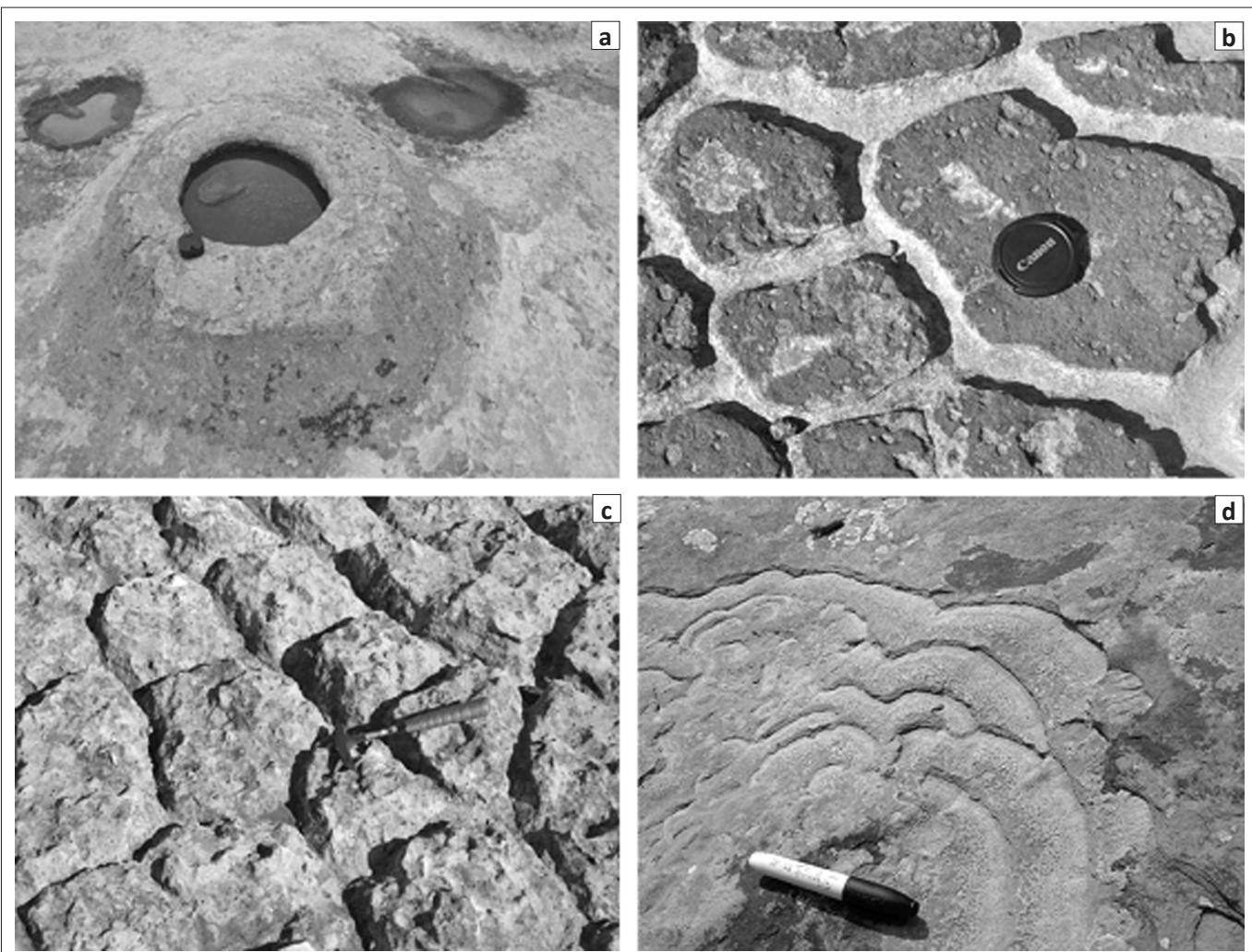


FIGURE 5: (a) A rock doughnut, (b) polygonal 'boxwork' structures, (c) convex polygonal structures and (d) lichen etching into Clarens Formation sandstone in the Golden Gate Highlands National Park.

5 cm – 50 cm high and have pit diameters ranging from 15 cm to 60 cm. Although reasons for the formation of rock doughnuts remain uncertain, Twidale and Campbell (1998) suggest that subsurface water held in the regolith was preferentially drained from zones adjacent to basins (doughnuts). The drained areas consequently had less moisture and thus weathered less rapidly than adjacent areas where more water was retained.

Polygonal cracking (tessellation or alligator cracking)

Polygonal fracture patterns consist mainly of pentagonal or hexagonal cracks meeting at ca 120° angles where they make tri-radial junctions on predominantly curved rock surface pavements (Robinson & Williams 1989, 1992; Young *et al.* 2009). Polygonal crack diameters may vary from 5 cm to more than 50 cm and in some instances micropolygonal cracking has been observed within larger tessellation plates (Robinson & Williams 1989). Although understanding the origin of tessellation has caused some debate (Young *et al.* 2009) and may differ between sites, it is a particularly common phenomenon in case-hardened rocks associated with regions of seasonal precipitation patterns (Robinson & Williams 1989, 1992; Twidale & Campbell 1993). Some of the likely causative mechanisms suggested include shrinkage of silica gel due to changing thermal and/or moisture conditions on the rock surface (Robinson & Williams 1989, 1992). These include thermal changes that set up differential surface stress (Branagan 1983; Croll 2009), expansion to accommodate precipitating solutions of iron, and thermal stresses or geochemical changes caused by emplacement of dolerite dykes (Velázquez *et al.* 2008).

Polygonal cracking is a common phenomenon in the GGHNP and in most instances occurs on horizontal case-hardened rock pavements, but may also be observed on vertical exposures such as at the base of cliffs along the Echo Ravine hiking trail and at Baboon Ridge. As observed in the field, the cracks may penetrate beyond the ca 5-mm thick case-hardened surface and continue further into the subsurface. At the Oribi Drive vulture site and also at Baboon Ridge, well-developed pavements occur in sandstones close to dolerite dykes. Less common are polygonal 'boxwork' structures where the joints are positive features, possibly due to silica induration along the joints, while the polygonal cells have been preferentially eroded to form small rock basins (Figure 5b).

Bedrock drainage microchannels (fluting, gutters, karren forms)

Bedrock drainage microchannels most frequently develop on limestone, dolomite and other soluble rock types. They are, however, also known to form on basaltic, granitic and sandstone surfaces as a result of sheet or channelised water flow across bedrock surfaces (forming rillenkarren or 'solution flutes' and rinnenkarren or rundkarren, respectively). The formation of such structures is primarily due to dissolution and biokarstic processes and/or the pre-existence of joints or surface irregularities (Crowther 1997; Dunkerley 1979; Fiol, Fornós & Ginés 1996; Wray 1997). In some instances long drainage microchannels are interrupted by shallow rock basins. Sandstone surfaces with ferruginous concretions are

common to the GGHNP and often lead to intricate surface drainage across bedrock, producing sharp-crested and highly fragmented surfaces. Trittkarren are much rarer and lesser-known weathering and erosion surface forms and have been described as 'small arcuate impressions' on sloping bedrock (Vincent 1983:205). Although examples of trittkarren have been observed in the GGHNP, their precise process origins remain poorly understood. However, Bögli (1960) suggested that such forms may develop where solution rates are influenced by the slow dissolution of gaseous CO₂ into the surface water in contact with bedrock and subsequently, when the microforms are established, they initiate self-sustaining processes such as concentrated overland flow towards the centre of each riser lip (Bögli 1980).

Twidale (1980) first described incised channelling into sandstone in the Drakensberg and suggested it to be the product of subsurface flow beneath a soil cover before the sandstone had been exposed. However, it is now more generally recognised that such sandstone channels (up to 100 mm wide and 10 mm deep) represent selective channelised drainage routes across rock surfaces (Robinson & Williams 1992; Wray 1997). Polygonally cracked surfaces permit drainage along the cracks, consequently accelerating erosion and eventually leading to well-incised bedrock microchannels, which grow in diameter and depth in the downflow direction. Observations at a site in the GGHNP indicate that originally fairly flat polygonally cracked surfaces may evolve into a more dramatically undulating micro-relief of incised channels (mean depth = 7 cm; *n* = 20) that separate broad convex polygonal tops with an average diameter of 33 cm (*n* = 40) (Figure 5c).

Lichens and weathering

Although it has previously been argued that lichens may protect rocks from enhanced weathering (Danin & Garty 1983), the more common view is that lichens promote rock weathering through chemical and mechanical processes which include metallic cation chelation, dissolution, swelling of organic and inorganic salts, and hyphal penetration, where expansion and contraction cause the detachment, incorporation and expulsion of rock flakes (Chen *et al.* 2000; McCarroll & Viles 1995). Wessels *et al.* (1995) have measured high strain on Clarens Formation sandstone surfaces in the GGHNP in association with increasing moisture uptake by some endolithic lichens, particularly at the thallus rim. In another study undertaken in the GGHNP, Wessels and Wessels (1991) identified bagworm larvae constructing their bags around lichen-weathered quartz crystals and thus demonstrated that lichen-induced weathering may operate synergistically with other biogenic processes on rock surfaces.

Studies, including some from the GGHNP, have indicated that some lichens may form circular or crescentic depressions bounded by microscars on rock surfaces (Wessels & Schoeman 1988; Wessels *et al.* 1995) (Figure 5d) owing to the combined actions of chemical weathering by chelation and mechanical disintegration associated with the penetration and expansion of hyphae (Cooks & Otto 1990; McKechnie, Grab & Drennan 2007; Moses & Smith 1993). Many of the flat

or gently sloping sandstone pavements in the GGHNP are extensively covered with circular patches of endolithic lichen thalli. When these thalli merge, they can produce remarkable flower-like patterns of pits in the rock surface (Figure 5d). Broadly similar forms have been recorded from the High Atlas Mountains in Morocco (Edwards *et al.* 2002). A survey undertaken on a rock wall in the GGHNP found that up to 29% of the colonisable surface was covered by thalli (Wessels & Schoeman 1988). Our own surveys supported this observation. The associated thalli microdepressions observed in the GGHNP include compound (i.e. microscars), coalescent and crescentic forms.

Lichenometric dating has become a common method for determining the relative age of exposed rock surfaces and has assisted in providing indications of glacial advances (Bradwell 2004), the timing of debris flows (Innes 1983), rates of fluvial downcutting (Thompson & Jones 1986) and past rockfall events (Bull 1996), to name a few. Surface age is traditionally determined by the single largest lichen, but other methods such as using the gradient of the lichen size-frequency distribution have also been used to produce well-calibrated dating curves (Bradwell 2004). Wessels and Schoeman (1988) examined freshly quarried Clarens Formation sandstone blocks used in the construction of a stone wall in 1942 to determine rates of lichen-induced weathering in the GGHNP. Their results suggest a theoretical mean surface denudation rate of 9.6 mm per 100 years where endolithic lichens are active on rock surfaces. Observations at the same site (the Glen Reenen graveyard) in 2009 showed that such lichen attack still continues. We measured 29 lichens (*Lecidea* aff. *sarcogynoides*) on 16 blocks and found that the average width of the lichens was 98.28 mm and that they penetrated into the rock surface to an average maximum depth of 4.8 mm. The average maximum rate of lichen denudation in the 67 years since 1942 was thus 7.16 mm per 100 years. Such results may prove valuable for any future attempt to determine the relative lichenometric-derived ages of mass movement phenomena in the GGHNP.

The depths of lichen denudation at Glen Reenen are not unique. For example, at a Clarens Formation sandstone at Baboon Ridge we measured 30 *Lecidea* pits and found that the average maximum width was 127 mm and the average maximum depth was 5.8 mm. At Red Mountain we measured 20 *Lecidea* pits in Elliot Formation sandstone and found that the comparable values were 65.25 mm and 3.3 mm, respectively.

Discussion and conclusion

Why does the GGHNP have such a rich diversity of macro- and microscale sandstone phenomena? Warke, McKinley and Smith (2006) have experimentally demonstrated that even minor structural and mineralogical differences between sandstone samples can result in substantially different weathering behaviours and consequent rock morphological outcomes. The GGHNP and foothills to the east of the Great Escarpment along KwaZulu-Natal and the Eastern Cape host a particularly dynamic assortment of sandstones, ranging from fine-grained mudstones and siltstones to

coarse-grained quartzitic sandstones. Given the relative abundance of feldspathic sandstones in the GGHNP, many of the macro- and microscale phenomena observed resemble those commonly recorded in granitic terrain. The presence of ferruginous chloritic mud and calcitic cement (Eriksson 1981) has provided the catalyst to the wide assortment of landforms associated with dissolution, which more typically occur in limestone or dolomitic terrain. Further, the rich diversity of the sandstone depositional history and concurrent neotectonics have provided for an equally complex spatial diversity of sandstone porosity, degree of lithification, and bedding and jointing characteristics, all of which contribute to sedimentary constructions and destructions. Finally, neither the past nor present climate can be underestimated for its all-important role in developing geomorphic phenomena through weathering, erosion and deposition of earth materials. The larger landforms (i.e. cliffs, pillars, arches, etc.) discussed in this paper are the likely product of past geological processes and controls and climate-driven geomorphic processes operating at geological timescales, while the smaller microforms (i.e. honeycombs, rock basins, polygonal cracks, etc.) are likely consequences of more recent climates (i.e. from the last few thousand years), characterised in particular by strong seasonal temperature and moisture differences.

Considerable work has focused on the deterioration of sandstone cultural rock, particularly in the context of monuments that belong to world cultural heritage (Mohammadi & Krumbein 2008) and sandstone buildings (Viles 2002). The rapid deterioration of sandstone containing rock art both locally (Hall, Meiklejohn & Arocena 2007; Meiklejohn 1997) and internationally (Benito *et al.* 1993; Walderhaug 1998) has also been a matter of concern in the GGHNP. The primary geomorphic causes for rock art deterioration are granular disintegration and flaking associated with salt crystallisation, hydration, wetting and drying, and chemical alterations (Benito *et al.* 1993; Meiklejohn 1997; Meiklejohn, Hall & Davis 2009). However, it should be recognised that for the most part, rock weathering is a natural process, and as Pope, Meierding and Paradise (2002:224) convey: 'geomorphic processes are integral in the creation of the "ruined" aesthetic'. Notwithstanding this, rock art requires preservation against unnecessary accelerated weathering, particularly where anthropogenic contributing factors are part of the cause.

Finally, an improved understanding of the mechanisms producing terrestrial sandstone phenomena, such as those described for the GGHNP, offers opportunities to better understand sedimentary and weathering processes on other planets such as Mars (Chan *et al.* 2008). Some of the common sandstone phenomena identified in the GGHNP and on Mars include surface flaking (exfoliation), case hardening, honeycombs, tafoni, 'blueberries' and polygonal cracking. The water-based mechanism for blueberry formation on Earth has been used as a terrestrial analogue for Martian blueberry formation and, more importantly, to provide possible evidence for the past or present occurrence of fluids on Mars (Squyres & Knoll 2005).

We hope that this work has extended scientific knowledge on sandstone phenomena in the GGHNP and thus provides a contribution to geoconservation and the regional geotourism industry through the provision of educational materials.

Acknowledgements

We thank South African National Parks for permission to carry out research at the GGHNP. Particular thanks to the manager of the GGHNP, Johan Taljaard, for his kind assistance and permission to use the photographic image of the natural stone bridge. S.G. acknowledges financial support from the Wits University Research Committee. Wendy Job kindly assisted in the cartographic production. We also appreciate the valuable comments and suggestions by the referees.

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